Modelling spatial and temporal dynamics of gross primary production in the Sahel from earth-observation-based photosynthetic capacity and quantum efficiency

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Abstract. It has been shown that vegetation growth in semi-arid regions is important to the global terrestrial CO₂ sink, which indicates the strong need for improved understanding and spatially explicit estimates of CO₂ uptake (gross primary production; GPP) in semi-arid ecosystems. This study has three aims: (1) to evaluate the MOD17A2H GPP (collection 6) product against GPP based on eddy covariance (EC) for six sites across the Sahel; (2) to characterize relationships between spatial and temporal variability in EC-based photosynthetic capacity ($F_{\text{opt}}$) and quantum efficiency ($\alpha$) and vegetation indices based on earth observation (EO) (normalized difference vegetation index (NDVI), renormalized difference vegetation index (RDVI), enhanced vegetation index (EVI) and shortwave infrared water stress index (SIWSI)); and (3) to study the applicability of EO upscaled $F_{\text{opt}}$ and $\alpha$ for GPP modelling purposes. MOD17A2H GPP (collection 6) drastically underestimated GPP, most likely because maximum light use efficiency is set too low for semi-arid ecosystems in the MODIS algorithm. Intra-annual dynamics in $F_{\text{opt}}$ were closely related to SIWSI being sensitive to equivalent water thickness, whereas $\alpha$ was closely related to RDVI being affected by chlorophyll abundance. Spatial and inter-annual dynamics in $F_{\text{opt}}$ and $\alpha$ were closely coupled to NDVI and RDVI, respectively. Modelled GPP based on $F_{\text{opt}}$ and $\alpha$ upscaled using EO-based indices reproduced in situ GPP well for all except a cropped site that was strongly impacted by anthropogenic land use. Upscaled GPP for the Sahel 2001–2014 was $736 \pm 39 \text{ g C m}^{-2} \text{ yr}^{-1}$. This study indicates the strong applicability of EO as a tool for spatially explicit estimates of GPP, $F_{\text{opt}}$ and $\alpha$; incorporating EO-based $F_{\text{opt}}$ and $\alpha$ in dynamic global vegetation models could improve estimates of vegetation production and simulations of ecosystem processes and hydro-biochemical cycles.

1 Introduction

Vegetation growth in semi-arid regions is an important sink for fossil fuel emissions. Mean carbon dioxide (CO₂) uptake by terrestrial ecosystems is dominated by highly productive lands, mainly tropical forests, whereas semi-arid regions are the main biome driving its inter-annual variability (Ahlström et al., 2015; Poulter et al., 2014). Semi-arid regions contribute to 60% of the long-term trend in the global terrestrial C sink (Ahlström et al., 2015). It is thus important to understand long-term variability of vegetation growth in semi-arid areas and the response of vegetation to environmental conditions to better quantify and forecast effects of climate change.

The Sahel is a semi-arid transition zone between the dry Sahara desert in the north and the humid Sudanian savanna in the south. The region has experienced numerous severe
droughts over the last decades, which resulted in region-wide famines in 1972–1973 and 1984–1985 and localized food shortages across the region in 1990, 2002, 2004, 2011 and 2012 (Abdi et al., 2014; United Nations, 2013). Vegetation production is thereby an important ecosystem service for livelihoods in the Sahel, but it is under threat. The region is experiencing strong population growth, increasing the demand on ecosystem services due to cropland expansion, increased pasture stocking rates and fuelwood extraction (Abdi et al., 2014).

At the same time as we have reports of declining vegetation production, we have contradicting reports of the greening of the Sahel based on earth observation (EO) data (Dardel et al., 2014; Fensholt et al., 2013). The greening of the Sahel has mainly been attributed to alleviated drought stress conditions due to increased precipitation since the mid-1990s (Hickler et al., 2005). Climate is thus another important factor regulating vegetation production. Semi-arid regions, such as the Sahel, are particularly vulnerable to climate fluctuations due to their dependency on moisture.

Estimation of gross primary production (GPP), i.e. uptake of atmospheric CO2 by vegetation, is still a major challenge for the remote sensing of ecosystem services. Gross primary production is a main driver of ecosystem services such as climate regulation, carbon (C) sequestration, C storage, food production and livestock grassland production. Within EO, spatial quantification of GPP generally involves light use efficiency (LUE), defined as the conversion efficiency of absorbed solar light into CO2 uptake (Monteith, 1972, 1977). It has been shown that LUE varies in space and time due to factors such as plant functional type, drought and temperature, nutrient levels, and physiological limitations of photosynthesis (Garbulsky et al., 2010; Paruelo et al., 2004; Kergoat et al., 2008). The LUE concept has been applied through various methods, either by using a biome-specific LUE constant (Ruimy et al., 1994) or by modifying a maximum LUE using meteorological variables (Running et al., 2004).

An example of a LUE-based model is the standard GPP product from the Moderate Resolution Imaging Spectroradiometer (MODIS) sensor (MOD17A2). Within the model, absorbed photosynthetically active radiation (PAR) is estimated as a product of the fraction of PAR absorbed by green vegetation (FPAR from MOD15A2) multiplied with daily PAR from the meteorological data of the Global Modeling and Assimilation Office (GMAO). A set of maximum LUE parameters specified for each biome are extracted from a Biome Properties Look-Up Table (BPLUT). Then maximum LUE is modified depending on air temperature (Tair) and vapour pressure deficit (VPD; Running et al., 2004). Sjöström et al. (2013) evaluated the MOD17A2 product (collection 5.1) for Africa and showed that it underestimated GPP for semi-arid savannas in the Sahel. Explanations for this underestimation were that the assigned maximum LUE from BPLUT was set too low and that there were uncertainties in the FPAR product (MOD15A2). Recently, a new collection of MOD17A2 at a 500 m spatial resolution was released (MOD17A2H, collection 6) with an updated BPLUT, updated GMAO meteorological data, improved quality control and gap-filling of the FPAR data from MOD15A2 (Running and Zhao, 2015).

It has been shown that the LUE method does not perform well in arid conditions and at agricultural sites (Turner et al., 2005). Additionally, the linearity assumed by the LUE model is not usually found as the response of GPP to incoming light follows more of an asymptotic curve (Cannell and Thornley, 1998). Investigating other methods for remotely determining GPP is thus of great importance, especially for semi-arid environments. Therefore, instead of LUE, we focus on the light response function of GPP at the canopy scale, and spatial and temporal variation of its two main parameters: maximum GPP under light saturation (canopy-scale photosynthetic capacity, Fopt) and the initial slope of the light response function (canopy-scale quantum efficiency, α; Falge et al., 2001; Tagesson et al., 2015a). Photosynthetic capacity is a measure of the maximum rate at which the canopy can fix CO2 during photosynthesis (µmol CO2 m⁻² s⁻¹), whereas α is the amount of CO2 fixed per incoming PAR (µmol CO2 µmol PAR⁻¹). To clarify the difference in LUE and α in this study, LUE (µmol CO2 µmol APAR⁻¹) is the slope of a linear fit between CO2 uptake and absorbed PAR, whereas α (µmol CO2 µmol PAR⁻¹) is the initial slope of an asymptotic curve against incoming PAR.

It has been proven that Fopt and α are closely related to chlorophyll abundance due to their coupling with the electron transport rate (Ide et al., 2010). Additionally, in semi-arid ecosystems, water availability is generally considered to be the main limiting factor affecting intra-annual dynamics of vegetation growth (Fensholt et al., 2013; Hickler et al., 2005; Tagesson et al., 2015b). Several remote sensing studies have established relationships between remotely sensed vegetation indices and ecosystem properties such as chlorophyll abundance and equivalent water thickness (Yoder and Pettigrew-Crosby, 1995; Fensholt and Sandholt, 2003). In this study, we will analyse whether EO vegetation indices can be used to upscale Fopt and α and investigate whether this could offer a promising way to map GPP in semi-arid areas. This potential will be analysed by the use of detailed ground observations from six eddy covariance (EC) flux tower sites across the Sahel.

The three aims of this study are

1. to investigate whether the recently released MOD17A2H GPP (collection 6) product is better at capturing GPP for the Sahel than collection 5.1. We hypothesize that the MOD17A2H GPP (collection 6) product will estimate GPP well for the six Sahelian EC sites because of major changes made in comparison to collection 5.1 (Running and Zhao, 2015);

2. to characterize the relationships between spatial and temporal variability in Fopt and α and remotely sensed...
vegetation indices. We hypothesise that EO vegetation indices that are closely related to chlorophyll abundance will be most strongly coupled with spatial and interannual dynamics in \( F_{\text{opt}} \) and \( \alpha \), whereas vegetation indices closely related to equivalent water thickness will be most strongly coupled with intra-annual dynamics in \( F_{\text{opt}} \) and \( \alpha \) across the Sahel;

3. to evaluate the applicability of a GPP model based on the light response function using EO vegetation indices and incoming PAR as input data.

2 Materials and methods

2.1 Site description

The Sahel stretches from the Atlantic Ocean in the west to the Red Sea in the east. The northern border towards the Sahara and the southern border towards the humid Sudanian savanna are defined by the 150 and 700 mm isohyets, respectively (Fig. 1; Prince et al., 1995). Tree and shrub canopy cover is now generally low (< 5 %) and dominated by species of Balanites, Acacia, Boscia and Combretaceae (Rietkerk et al., 1996). Annual grasses such as Schoenefeldia gracilis, Dactyloctenium aegyptium, Aristida mutabilis and Cenchrus biflorus dominate the herbaceous layer, but perennial grasses such as Andropogon gayanus and Cymbopogon schoenanthus can also be found (Rietkerk et al., 1996; de Ridder et al., 1982). From the FLUXNET database (Baldocchi et al., 2001) we selected six measurement sites with EC-based CO\(_2\) flux data from the Sahel (Table 1; Fig. 1). The sites represent a variety of ecosystems present in the region, from dry fallow bush savanna to seasonally inundated acacia forest. For a full description of the measurement sites, we refer to Tagesson et al. (2016a) and references in Table 1.

2.2 Data collection

2.2.1 Eddy covariance and hydrometeorological in situ data

Eddy covariance and hydrometeorological data originating from the years between 2005 and 2013 were collected from the principal investigators of the measurement sites (Tagesson et al., 2016a). The EC sensor set-up consisted of open-path CO\(_2\) and H\(_2\)O infrared gas analysers and three-axis sonic anemometers. Data were collected at 20 Hz and statistics were calculated for 30 min periods. For a full description of the sensor set-up and post processing of EC data, see the references in Table 1. Final fluxes were filtered according to quality flags provided by FLUXNET and outliers were filtered according to Papale et al. (2006). We extracted the original net ecosystem exchange (NEE) data without any gap-filling or partitioning of NEE to GPP and ecosystem respiration. The hydrometeorological data collected were air temperature \( (T_{\text{air}}, ^{\circ} \text{C}) \), rainfall \( (P, \text{mm}) \), relative air humidity \( (\text{Rh}, \%) \), soil moisture at 0.1 m depth \( (\text{SWC}, \% \text{volumetric water content}) \), incoming global radiation \( (R_g, \text{W m}^{-2}) \), incoming photosynthetically active radiation \( \text{(PAR, } \mu\text{mol m}^{-2} \text{s}^{-1}) \), VPD \( (\text{hPa}) \), peak dry weight biomass \( (\text{g dry weight m}^{-2}) \), C3 / C4 species ratio and soil conditions \( (\text{nitrogen and C concentration, } \%) \). For a full description of the collected data and sensor set-up, see Tagesson et al. (2016a).

2.2.2 Earth observation data and gridded ancillary data

Composite products from MODIS/Terra covering the Sahel were acquired at Reverb ECHO (NASA, 2016). Collected products were GPP (MOD17A2H, collection 6), nadir bidirectional reflectance distribution function adjusted reflectance \( \text{(NBAR; 8-day composites; MCD43A4, collection 5.1) at } 500 \times 500 \text{ m}^{2} \text{ spatial resolution, the normalized difference vegetation index (NDVI) and the enhanced vegetation index (EVI; 16-day composites; MOD13Q1; collection 6) at } 250 \times 250 \text{ m}^{2} \text{ spatial resolution. The NBAR product was preferred over the reflectance product (MOD09A1) in order to avoid variability caused by varying sun and sensor viewing geometry (Huber et al., 2014; Tagesson et al., 2015c). We extracted the median of 3 × 3 pixels centred at the location of each EC tower. Time series of EO products were filtered according to MODIS quality control data. MOD17A2H is a gap-filled and filtered product, and QC data from MCD43A2 were used for filtering of MCD43A4. Bit 2–5 (highest-decreasing quality) was used for MOD13Q1. Finally, data were gap-filled to daily values using linear interpolation.

We downloaded ERA Interim reanalysis PAR at the ground surface \( \text{(W m}^{-2}) \) with a spatial resolution of \( 0.25^\circ \times 0.25^\circ \text{ accumulated for each 3 h period from 2000 to 2014 from the European Centre for Medium-Range Weather Forecasts (ECMWF) (Dee et al., 2011; ECMWF, 2016a).}

2.3 Data handling

2.3.1 Intra-annual dynamics in photosynthetic capacity and quantum efficiency

To estimate daily values of EC-based \( F_{\text{opt}} \) and \( \alpha \), the asymptotic Mitscherlich light-response function was fitted between daytime NEE and incoming PAR using a 7-day moving window with a 1-day time step:

\[
\text{NEE} = -(F_{\text{opt}}) \times (1 - e^{-\alpha \times \text{PAR} / F_{\text{opt}}}) + R_d.
\]

where \( F_{\text{opt}} \) is CO\(_2\) uptake at light saturation (photosynthetic capacity; \( \mu\text{mol CO}_2 \text{m}^{-2} \text{s}^{-1} \)), \( R_d \) is dark respiration \( (\mu\text{mol CO}_2 \text{m}^{-2} \text{s}^{-1}) \) and \( \alpha \) is the initial slope of the light response curve (quantum efficiency; \( \mu\text{mol CO}_2 \mu\text{mol PAR}^{-1} \)).
Table 1. Description of the six measurement sites, including location, soil type, ecosystem type and dominant species.

<table>
<thead>
<tr>
<th>Measurement site</th>
<th>Coordinates</th>
<th>Soil type</th>
<th>Ecosystem</th>
<th>Dominant species</th>
</tr>
</thead>
<tbody>
<tr>
<td>Agoufou&lt;sup&gt;a&lt;/sup&gt;</td>
<td>15.34° N, 1.48° W</td>
<td>Sandy ferruginous Arenosol</td>
<td>Open woody savanna (4 % tree cover)</td>
<td>Trees: Acacia spp., Balanites aegyptiaca, Combretum glutinosum, Herbs: Zornia glochidiata, Cenchrus biflorus, Aristida mutabilis, Tragus berteronianus</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td>Dahra&lt;sup&gt;b&lt;/sup&gt;</td>
</tr>
<tr>
<td>(ML-AgG, Mali)</td>
<td></td>
<td></td>
<td></td>
<td>15.40° N, 15.43° W Sandy luvic Arenosol Grassland and/or shrubland Savanna (3 % tree cover)</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td>Demokeya&lt;sup&gt;c&lt;/sup&gt;</td>
</tr>
<tr>
<td>(SD-Dem, Sudan)</td>
<td>13.28° N, 30.48° E</td>
<td>Cambic Arenosol</td>
<td>Sparse acacia savannah (7 % tree cover)</td>
<td>Trees: Acacia seyal, Acacia nilotica, Balanites aegyptiaca, Herbs: Sporobolus revolvens, Echinocloa colona, Aeschynomene senistive</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td>Kelma&lt;sup&gt;a&lt;/sup&gt;</td>
</tr>
<tr>
<td>(ML-Kem, Mali)</td>
<td>15.22° N, 1.57° W</td>
<td>Clay soil depression</td>
<td>Open acacia forest (90 % tree cover)</td>
<td>Wankama Fallow&lt;sup&gt;d&lt;/sup&gt;</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td>13.65° N, 2.63° E Sandy ferruginous Arenosol Fallow bush</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td>Wankama Millet&lt;sup&gt;e&lt;/sup&gt;</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td></td>
<td>13.64° N, 2.63° E Sandy ferruginous Arenosol Millet crop</td>
</tr>
</tbody>
</table>

Figure 1. Land cover classes for the Sahel and the location of the six measurement sites of this study. The land cover classes are based on multi-sensor satellite observations (Mayaux et al., 2003). The sites are Agoufou (ML-AgG), Dahra (SN-Dah), Demokeya (SD-Dem), Kelma (ML-Kem), Wankama Fallow (NE-WaF) and Wankama Millet (NE-WaM). The thick black line delineates borders of the Sahel based on annual 150 and 700 mm isohyets (Prince et al., 1995).

<sup>a</sup> Timouk et al. (2009). <sup>b</sup> Tagesson et al. (2015b). <sup>c</sup> Sjöström et al. (2009). <sup>d</sup> Velluet et al. (2014). <sup>e</sup> Boulain et al. (2009).
Falge et al., 2001). By subtracting $R_d$ from Eq. (1), the function was forced through zero and GPP was thereby estimated. To ensure a high quality of fitted parameters, parameters were excluded from the analysis when fitting was insignificant ($p$ value $>0.05$) and when they were out of range ($F_{\text{opt}}$ and $\alpha$ > peak value of the rainy season times 1.2). Additionally, outliers were filtered following the method by Papale et al. (2006) using a 30-day moving window with a 1-day time step.

### 2.3.2 Vegetation indices

The maximum absorption in red wavelengths generally occurs at 682 nm as this is the peak absorption for chlorophyll $a$ and $b$ (Thenkabail et al., 2000), which makes vegetation indices that include the red band sensitive to chlorophyll abundance. By far the most common vegetation index is NDVI (Rouse et al., 1974):

$$\text{NDVI} = \frac{(\rho_{\text{NIR}} - \rho_{\text{red}})}{(\rho_{\text{NIR}} + \rho_{\text{red}})},$$

where $\rho_{\text{NIR}}$ is the reflectance factor in the near-infrared (NIR) band (band 2) and $\rho_{\text{red}}$ is the reflectance factor in the red band (band 1). Near-infrared radiance is reflected by leaf cells since absorption of these wavelengths would result in overheating of the plant, whereas red radiance is absorbed by chlorophyll and its accessory pigments (Gates et al., 1965). Normalization is done to reduce effects of atmospheric errors, solar zenith angles and sensor viewing geometry, as well as to increase the vegetation signal (Qi et al., 1994; Inoue et al., 2008).

A well-known deficiency of NDVI is problems of index saturation at high biomass because absorption of red light at $\sim 670$ nm peaks at higher biomass loads, whereas NIR reflectance continues to increase due to multiple scattering effects (Mutanga and Skidmore, 2004; Jin and Eklundh, 2014). By reducing atmospheric and soil background influences, EVI is designed to increase the signal from the vegetation and maintain sensitivity in high biomass regions (Huete et al., 2002).

$$\text{EVI} = G \frac{(\rho_{\text{NIR}} - \rho_{\text{red}})}{(\rho_{\text{NIR}} + C_1 \rho_{\text{red}} - C_2 \rho_{\text{blue}} + L)},$$

where $\rho_{\text{blue}}$ is the reflectance factor in the blue band (band 3). The coefficients $C_1 = 6$ and $C_2 = 7.5$ correct for atmospheric influences, while $L = 1$ adjusts for the canopy background. The factor $G = 2.5$ is a gain factor.

Another attempt to overcome problems of NDVI saturation was proposed by Roujean and Breon (1995), who suggested the renormalized difference vegetation index (RDVI), which combines advantages of DVI (NIR-red) and NDVI for low and high vegetation cover, respectively:

$$\text{RDVI} = \frac{(\rho_{\text{NIR}} - \rho_{\text{red}})}{\sqrt{(\rho_{\text{NIR}} + \rho_{\text{red}})}}.$$
indices for all sites combined were also analysed. In the analysis for all sites, data were normalized to avoid influence of spatial and inter-annual variability. Time series of ratios of \( F_{\text{opt}} \) and \( \alpha \) (\( F_{\text{opt frac}} \) and \( \alpha_{\text{frac}} \)) against the annual peak values (\( F_{\text{opt peak}} \) and \( \alpha_{\text{peak}} \); see below for calculation of annual peak values) were estimated for all sites:

\[
F_{\text{opt frac}} = \frac{F_{\text{opt}}}{F_{\text{opt peak}}}, \quad \alpha_{\text{frac}} = \frac{\alpha}{\alpha_{\text{peak}}}. \quad (7)
\]

The same standardization procedure was used for all vegetation indices (\( V_{\text{frac}} \)):

\[
V_{\text{frac}} = \frac{V_{\text{peak}}}{V_{\text{frac peak}}}, \quad (9)
\]

where \( V_{\text{peak}} \) is the annual peak values of the vegetation indices (14-day running mean with highest annual value). The \( \alpha_{\text{frac}} \) and \( F_{\text{opt frac}} \) were correlated with the different \( V_{\text{frac}} \) to investigate the coupling between intra-annual dynamics in \( F_{\text{opt}} \) and \( \alpha \) and the vegetation indices for all sites.

Regression trees were used to fill gaps in the daily estimates of \( F_{\text{opt}} \) and \( \alpha \). Based on 100 cross-validation runs, 100 tree sizes were chosen, and these trees were then used to estimate \( F_{\text{opt}} \) and \( \alpha \) following the method in De’ath and Fabricius (2000). We used SWC, VPD, \( T_{\text{air}} \), PAR and the vegetation index with the strongest correlation with intra-annual dynamics as explanatory variables in the analysis. In the analysis for all sites, the same standardization procedure as done for \( F_{\text{opt}} \) and \( \alpha \) and the vegetation indices was done for the hydrometeorological variables. The 100 \( F_{\text{opt}} \) and \( \alpha \) output subsets from the regression trees were averaged and used for filling gaps in the time series of \( F_{\text{opt}} \) and \( \alpha \). From these time series, we estimated annual peak values of \( F_{\text{opt}} \) and \( \alpha \) (\( F_{\text{opt peak}} \) and \( \alpha_{\text{peak}} \)) as the 14-day running mean with the highest annual value. To investigate spatial and inter-annual variability in \( F_{\text{opt}} \) and \( \alpha \) across the measurement sites of the Sahel, \( F_{\text{opt peak}} \) and \( \alpha_{\text{peak}} \) were correlated with the annual sum of \( P \); yearly means of \( T_{\text{air}} \), SWC, RH, VPD and \( R_{\text{g}} \); annual peak values of biomass; soil nitrogen and C concentrations; the C3/C4 ratio; and \( V_{\text{frac peak}} \).

### 2.4.2 Parameterization and evaluation of the GPP model and evaluation of the MODIS GPP

On the basis of Eq. (1) and the outcome of the statistical analysis previously described under Sect. 2.4.1 (for results, see Sect. 3.2), a model for estimating GPP across the Sahel was created:

\[
\text{GPP} = -F_{\text{opt}} \times \left( 1 - e^{-\frac{\alpha \times \text{PAR}}{F_{\text{opt}}}} \right). \quad (10)
\]

Firstly, \( F_{\text{opt peak}} \) and \( \alpha_{\text{peak}} \) were estimated spatially and inter-annually using linear regression functions fitted against the vegetation indices, with the strongest relationships to spatial and inter-annual variability in \( F_{\text{opt}} \) and \( \alpha \) for all sites. Secondly, exponential regression functions were established for \( F_{\text{opt frac}} \) and \( \alpha_{\text{frac}} \) with the vegetation index, with the strongest relationships with intra-annual variability of \( F_{\text{opt frac}} \) and \( \alpha_{\text{frac}} \) for all sites. By combining these relationships, \( F_{\text{opt}} \) and \( \alpha \) can be calculated for any day of year and for any point in space across the Sahel:

\[
F_{\text{opt}} = F_{\text{opt peak}} \times F_{\text{opt frac}} = (k_{\text{Fopt}} \times \text{NDVI}_{\text{peak}} + m_{\text{Fopt}}) \\
\left( n_{\text{Fopt}} \times e^{(l_{\text{Fopt}} \times \text{RDVI}_{\text{frac}})} \right), \quad (11)
\]

\[
\alpha = \alpha_{\text{peak}} \times \alpha_{\text{frac}} = (k_{\alpha} \times \text{RDVI}_{\text{peak}} + m_{\alpha}) \\
\left( n_{\alpha} \times e^{(l_{\alpha} \times \text{RDVI}_{\text{frac}})} \right), \quad (12)
\]

where \( k_{\text{Fopt}} \) and \( k_{\alpha} \) are slopes and \( m_{\text{Fopt}} \) and \( m_{\alpha} \) are intercepts of the linear regressions, giving \( F_{\text{opt peak}} \) and \( \alpha_{\text{peak}} \), respectively; \( l_{\text{Fopt}} \) and \( l_{\alpha} \) are coefficients and \( n_{\text{Fopt}} \) and \( n_{\alpha} \) are intercepts of the exponential regressions, giving \( F_{\text{opt frac}} \) and \( \alpha_{\text{frac}} \), respectively. Equations (11) and (12) were inserted into Eq. (10), and GPP was thereby estimated as

\[
\text{GPP} = -\left( F_{\text{opt peak}} \times F_{\text{opt frac}} \right) \times \left( 1 - e^{\left( -\frac{\alpha_{\text{peak}} \times \text{PAR}}{F_{\text{opt peak}} \times F_{\text{opt frac}}} \right)} \right) = \\
- \left( k_{\text{Fopt}} \times \text{NDVI}_{\text{peak}} + m_{\text{Fopt}} \right) \left( n_{\text{Fopt}} \times e^{(l_{\text{Fopt}} \times \text{RDVI}_{\text{frac}})} \right) \times \left( 1 - e^{\left( -\frac{\alpha_{\text{peak}} \times \text{RDVI}_{\text{peak}} + m_{\alpha}}{n_{\text{Fopt}} \times e^{(l_{\text{Fopt}} \times \text{RDVI}_{\text{frac}})}} \right)} \right). \quad (13)
\]

A bootstrap simulation methodology was used when fitting the least squares regression functions for parameterization of the GPP model (Richter et al., 2012). For each of the iterations, some of the EC site years were included and some were omitted. The bootstrap simulations generated 200 sets of \( k_{\text{Fopt}}, k_{\alpha}, m_{\text{Fopt}}, m_{\alpha}, l_{\text{Fopt}}, l_{\alpha}, n_{\text{Fopt}}, n_{\alpha} \) and coefficient of determination (\( R^2 \)). Possible errors (e.g. random sampling errors, aerosols, electrical sensor noise, filtering and gap-filling errors, clouds and satellite sensor degradation) can be present in both the predictor and the response variables. Hence, we selected reduced major axis regressions to account for errors in both predictor and response variables when fitting the regression functions. The regression models were validated against the omission site years within the bootstrap simulation methodology by calculating the RMSE and by fitting an ordinate least square linear regression between modelled and independent variables.

Similarly, the MODIS GPP product (MOD17A2H; collection 6) was evaluated against independent GPP from the EC
sites by calculating the RMSE and by fitting an ordinary least square linear regression.

3 Results

3.1 Evaluation of the MODIS GPP product

There was a strong linear relationship between the MODIS GPP product (MOD17A2H; collection 6) and independent GPP (slope = 0.17; intercept = 0.11 g C m$^{-2}$ d$^{-1}$; $R^2 = 0.69$; $n = 598$). However, MOD17A2H strongly underestimated independent GPP, resulting in a high RMSE (2.69 g C m$^{-2}$ d$^{-1}$). It can be seen that some points for the Kelma site were quite low for MOD17A2H, whereas they were relatively high for the independent GPP (Fig. 2). Kelma is an inundated Acacia forest located in a clay soil depression. These differentiated values were found in the beginning of the dry season, when the depression was still inundated, whereas the larger area was turning dry.

3.2 Intra-annual dynamics in photosynthetic capacity and quantum efficiency

Intra-annual dynamics in $F_{\text{opt}}$ and $\alpha$ differed in amplitude, but were otherwise similar across the measurement sites in the Sahel (Fig. 3). There was no green ground vegetation during the dry season, and the low photosynthetic activity was due to few evergreen trees. This resulted in low values for both $F_{\text{opt}}$ and $\alpha$ during the dry season. The vegetation responded strongly to rainfall, and both $F_{\text{opt}}$ and $\alpha$ increased during the early phase of the rainy season. Generally, $F_{\text{opt}}$ peaked slightly earlier than $\alpha$ (average ±1 SD: 7 ± 10 days; Fig. 3).

All vegetation indices described intra-annual dynamics in $F_{\text{opt}}$ reasonably well at all sites (Table 2). The vegetation index SIWSI$_{12}$ had the highest correlation for all sites except Wankama Millet, where it was RDVI. When all sites were combined, all indices described seasonality in $F_{\text{opt}}$ well, but RDVI had the strongest correlation (Table 2).

Intra-annual dynamics in $\alpha$ were also closely coupled to intra-annual dynamics in the vegetation indices for all sites (Table 2). For $\alpha$, RDVI was the strongest index describing intra-annual dynamics, except for Wankama Fallow, where it was EVI. When all sites were combined, all indices described intra-annual dynamics in $\alpha$ well, but RDVI was still the index with the strongest relationship (Table 2).
The regression trees used for gap-filling explained the intra-annual dynamics in $F_{opt}$ and $\alpha$ well for all sites (Table 3; Fig. S2). The regression trees explained intra-annual dynamics in $F_{opt}$ better than in $\alpha$, and multi-year sites were better predicted than single-year sites (Fig. S2). The main explanatory variables coupled to intra-annual dynamics in $F_{opt}$ for all sites across the Sahel were on the order of RDVI, SWC, VPD, $T_{air}$, and PAR. For $\alpha$, they were RDVI, SWC, VPD and $T_{air}$ (Table 3). The strong relationship to SWC and VPD indicates drought stress during periods of low rainfall. For all sites across the Sahel, incorporating hydrometeorological variables increased the ability to determine intra-annual dynamics in $F_{opt}$ and $\alpha$ compared to the ordinary least square linear regressions against vegetation indices (Table 2, data given as $r$; Table 3; Figs. 3 and S2). For all sites, incorporation of these variables increased $R^2$ from 0.81 to 0.87 and from 0.74 to 0.84 for $F_{opt}$ and $\alpha$, respectively.

### 3.3 Spatial and inter-annual dynamics in photosynthetic capacity and quantum efficiency

Large spatial and inter-annual variability in $F_{opt\_peak}$ and $\alpha_{peak}$ were found across the six measurement sites; $F_{opt\_peak}$ ranged between 10.1 (Wankama Milliet 2005) and 50.0 $\mu$mol CO$_2$ m$^{-2}$ s$^{-1}$ (Dahra 2010), and $\alpha_{peak}$ ranged between 0.020 (Demokeya 2007) and 0.064 $\mu$mol CO$_2$ $\mu$mol PAR$^{-1}$ (Dahra 2010; Table 4). The average 2-week running mean peak values of $F_{opt}$ and $\alpha$ for all sites were 26.4 $\mu$mol CO$_2$ m$^{-2}$ s$^{-1}$ and 0.040 $\mu$mol CO$_2$ $\mu$mol PAR$^{-1}$, respectively. All vegetation
Table 4. Annual peak values of quantum efficiency ($\alpha_{\text{peak}}$; µmol CO$_2$ µmol PAR$^{-1}$) and photosynthetic capacity ($F_{\text{opt,peak}}$; µmol CO$_2$ m$^{-2}$ s$^{-1}$) for the six measurement sites (Fig. 1). The peak values are the 2-week running mean with highest annual value.

<table>
<thead>
<tr>
<th>Measurement site</th>
<th>Year</th>
<th>$\alpha_{\text{peak}}$</th>
<th>$F_{\text{opt,peak}}$</th>
</tr>
</thead>
<tbody>
<tr>
<td>ML-AgG</td>
<td>2007</td>
<td>0.0396</td>
<td>24.5</td>
</tr>
<tr>
<td>SN-Dah</td>
<td>2010</td>
<td>0.0638</td>
<td>50.0</td>
</tr>
<tr>
<td></td>
<td>2011</td>
<td>0.0507</td>
<td>42.3</td>
</tr>
<tr>
<td></td>
<td>2012</td>
<td>0.0480</td>
<td>39.2</td>
</tr>
<tr>
<td></td>
<td>2013</td>
<td>0.0549</td>
<td>40.0</td>
</tr>
<tr>
<td>SD-Dem</td>
<td>2007</td>
<td>0.0257</td>
<td>16.5</td>
</tr>
<tr>
<td></td>
<td>2008</td>
<td>0.0327</td>
<td>21.0</td>
</tr>
<tr>
<td></td>
<td>2009</td>
<td>0.0368</td>
<td>16.5</td>
</tr>
<tr>
<td>ML-Kem</td>
<td>2007</td>
<td>0.0526</td>
<td>33.5</td>
</tr>
<tr>
<td>NE-WaF</td>
<td>2005</td>
<td>0.0273</td>
<td>18.2</td>
</tr>
<tr>
<td></td>
<td>2006</td>
<td>0.0413</td>
<td>21.0</td>
</tr>
<tr>
<td>NE-WaM</td>
<td>2005</td>
<td>0.0252</td>
<td>10.6</td>
</tr>
<tr>
<td></td>
<td>2006</td>
<td>0.0200</td>
<td>10.1</td>
</tr>
<tr>
<td><strong>Average</strong></td>
<td></td>
<td><strong>0.0399</strong></td>
<td><strong>26.4</strong></td>
</tr>
</tbody>
</table>

indices determined spatial and inter-annual dynamics well in both $F_{\text{opt,peak}}$ and $\alpha_{\text{peak}}$ (Table 5); $F_{\text{opt,peak}}$ was most closely coupled with NDVI$_{\text{peak}}$, whereas $\alpha_{\text{peak}}$ was more closely coupled with RDVI$_{\text{peak}}$ (Fig. 4). $F_{\text{opt,peak}}$ also correlated well with peak dry weight biomass, C content in the soil and RH, whereas $\alpha_{\text{peak}}$ also correlated with peak dry weight biomass and C content in the soil (Table 5).

3.4 Spatially extrapolated photosynthetic capacity, quantum efficiency and gross primary production across the Sahel and evaluation of the GPP model

The spatially extrapolated $F_{\text{opt}}$, $\alpha$ and GPP averaged over the Sahel for 2001–2014 were 22.5 ± 1.7 µmol CO$_2$ m$^{-2}$ s$^{-1}$, 0.030 ± 0.002 µmol CO$_2$ µmol PAR$^{-1}$ and 736 ± 39 g C m$^{-2}$ yr$^{-1}$, respectively. At a regional scale, it can be seen that $F_{\text{opt}}$, $\alpha$ and GPP decreased substantially with latitude (Fig. 5). The highest values were found in southeastern Senegal, western Mali, in parts of southern Sudan and on the border between Sudan and South Sudan. The lowest values were found along the northernmost parts of the Sahel on the border to the Sahara in Mauritania, in northern Mali and in northern Niger.

Modelled GPP was similar to independent GPP on average, and there was a strong linear relationship between modelled GPP and independent GPP for all sites (Fig. 6; Table 6). However, when separating the evaluation between measurement sites, it can be seen that the model reproduced some sites better than others (Fig. 7; Table 6). Wankama Millet was generally overestimated, whereas on average the model worked well for Demokaya but underestimated high values (Fig. 7; Table 6). Variability of independent GPP at the other sites was reproduced by the model reasonably well (Fig. 7; Table 6). The final parameters of the GPP model (Eq. 13) are shown in Table 7.

4 Discussion

Our hypothesis that vegetation indices closely related to SIWSI would be most strongly coupled with intra-annual dynamics in $F_{\text{opt}}$ and $\alpha$ was not rejected for $F_{\text{opt}}$ since this was the case for all sites except for Wankama Millet (Table 2). However, our hypothesis was rejected for $\alpha$, since it was more closely related to vegetation indices of chlorophyll
Figure 4. Scatter plots of annual peak values for the six measurement sites (Fig. 1) of (a) photosynthetic capacity ($F_{\text{opt\_peak}}$) and (b) quantum efficiency ($\alpha_{\text{peak}}$) against peak values of normalized difference vegetation index (NDVI$_{\text{peak}}$) and renormalized difference vegetation index (RDVI$_{\text{peak}}$), respectively. The annual peak values were estimated by taking the annual maximum of a 2-week running mean.

Table 6. Statistics regarding the evaluation of the gross primary production (GPP) model for the six measurement sites (Fig. 1). In situ and modelled GPP are averages ± 1 SD. RMSE is the root mean square error, and slope, intercept and $R^2$ are from the fitted ordinary least square linear regressions.

<table>
<thead>
<tr>
<th>Measurement site</th>
<th>In situ GPP (µmol CO$_2$ m$^{-2}$ s$^{-1}$)</th>
<th>Modelled GPP (µmol CO$_2$ m$^{-2}$ s$^{-1}$)</th>
<th>RMSE (µmol CO$_2$ m$^{-2}$ s$^{-1}$)</th>
<th>Slope (µmol CO$_2$ m$^{-2}$ s$^{-1}$)</th>
<th>Intercept (µmol CO$_2$ m$^{-2}$ s$^{-1}$)</th>
<th>$R^2$</th>
</tr>
</thead>
<tbody>
<tr>
<td>ML-AgG</td>
<td>5.35 ± 6.38</td>
<td>5.97 ± 5.80</td>
<td>2.48 ± 0.10</td>
<td>0.84 ± 0.003</td>
<td>1.46 ± 0.01</td>
<td>0.86 ± 0.002</td>
</tr>
<tr>
<td>SN-Dah</td>
<td>9.39 ± 10.17</td>
<td>8.87 ± 9.67</td>
<td>3.99 ± 1.34</td>
<td>0.88 ± 0.002</td>
<td>0.62 ± 0.01</td>
<td>0.85 ± 0.001</td>
</tr>
<tr>
<td>SD-Dem</td>
<td>4.26 ± 4.55</td>
<td>3.98 ± 3.90</td>
<td>3.15 ± 1.06</td>
<td>0.63 ± 0.003</td>
<td>1.31 ± 0.007</td>
<td>0.54 ± 0.02</td>
</tr>
<tr>
<td>ML-Kem</td>
<td>11.16 ± 8.02</td>
<td>10.52 ± 9.22</td>
<td>4.35 ± 1.23</td>
<td>1.02 ± 0.003</td>
<td>−0.82 ± 0.03</td>
<td>0.78 ± 0.002</td>
</tr>
<tr>
<td>NE-WaF</td>
<td>5.77 ± 4.17</td>
<td>6.63 ± 3.53</td>
<td>2.47 ± 1.05</td>
<td>0.70 ± 0.005</td>
<td>2.58 ± 0.02</td>
<td>0.69 ± 0.003</td>
</tr>
<tr>
<td>NE-WaM</td>
<td>3.04 ± 1.93</td>
<td>6.35 ± 3.47</td>
<td>4.12 ± 0.99</td>
<td>1.31 ± 0.004</td>
<td>2.37 ± 0.02</td>
<td>0.53 ± 0.003</td>
</tr>
<tr>
<td>Average</td>
<td>6.73 ± 7.72</td>
<td>7.02 ± 7.39</td>
<td>3.68 ± 0.55</td>
<td>0.83 ± 0.07</td>
<td>1.34 ± 0.82</td>
<td>0.84 ± 0.07</td>
</tr>
</tbody>
</table>

Table 7. The parameters for Eq. (13) that were used in the final gross primary production (GPP) model. RMSE is the root mean square error, and $R^2$ is the coefficient of determination for the regression models predicting the different variables.

<table>
<thead>
<tr>
<th>Parameter</th>
<th>Value</th>
<th>RMSE</th>
<th>$R^2$</th>
</tr>
</thead>
<tbody>
<tr>
<td>$k_{F_{\text{opt}}}$</td>
<td>79.6 ± 6.3</td>
<td>5.1 ± 1.3</td>
<td>0.89 ± 0.05</td>
</tr>
<tr>
<td>$m_{F_{\text{opt}}}$</td>
<td>−7.3 ± 3.2</td>
<td>0.15 ± 0.02</td>
<td>0.88 ± 0.06</td>
</tr>
<tr>
<td>$l_{F_{\text{opt}}}$</td>
<td>3.51 ± 0.19</td>
<td>0.03 ± 0.006</td>
<td>0.16 ± 0.02</td>
</tr>
<tr>
<td>$n_{F_{\text{opt}}}$</td>
<td>0.02 ± 0.07</td>
<td>0.20 ± 0.02</td>
<td>0.80 ± 0.10</td>
</tr>
</tbody>
</table>

abundance (RDVI and EVI). In the Sahel, soil moisture conditions in the early rainy season are important for vegetation growth, and during this phase vegetation is especially vulnerable to drought conditions (Rockström and de Rouw, 1997; Tagesson et al., 2016a; Mbow et al., 2013). Photosynthetic capacity ($F_{\text{opt}}$) peaked earlier than $\alpha$ did in the rainy season (Fig. 3), thereby explaining the close relationship of $F_{\text{opt}}$ to SIWSI. Leaf area index increased over the growing season and leaf area index is closely coupled with vegetation indices related to chlorophyll abundance (Tagesson et al., 2009).
increase in leaf area index increased canopy level quantum efficiency (α), thereby explaining the closer relationship of α with RDVI.

Our hypothesis that vegetation indices closely related to chlorophyll abundance would be most strongly coupled with spatial and inter-annual dynamics in F_{opt} and α was not rejected for either F_{opt} or α; NDVI, EVI and RDVI all correlated with spatial and inter-annual dynamics in F_{opt} and α (Table 5). However, it was surprising that NDVI_{peak} had the strongest correlation with spatial and inter-annual variability in F_{opt} (Table 5). Both EVI and RDVI should be less sensitive to saturation effects than NDVI (Huete et al., 2002; Roujean and Breon, 1995), and based on this it can be assumed that peak values of these indices should have stronger relationships with peak values of F_{opt} and α. However, vegetation indices with a high sensitivity to changes in green biomass at high biomass loads become less sensitive to green biomass changes at low biomass loads (Huete et al., 2002). The peak leaf area index for ecosystems across the Sahel is generally ∼2 m² m⁻² or less, whereas the saturation issue of NDVI generally starts at a leaf area index of about 2–5 m² m⁻² (Haboudane et al., 2004).

The F_{opt_peak} estimates from Agoufou, Demokeya and the Wankama sites were similar, whereas Dahra and Kelma values were high in relation to previously reported canopy-scale F_{opt_peak} from the Sahel (∼−8 to −23 µmol m⁻² s⁻¹; Hanan et al., 1998; Merbold et al., 2009; Moncrieff et al., 1997; Boulain et al., 2009; Levy et al., 1997; Moneny et al., 1997). These previous studies reported much lower F_{opt} at canopy scale than at leaf scale (e.g. Levy et al., 1997: 10 vs. 44 µmol m⁻² s⁻¹; Boulain et al., 2009: 8 vs. 50 µmol m⁻² s⁻¹). The leaf area index at Dahra and Kelma peaked at 2.1 and 2.7, respectively (Timouk et al., 2009; Tagesson et al., 2015a), and it was substantially higher than at the above-mentioned sites. A possible explanation for high F_{opt} estimates at Dahra and Kelma could therefore be the higher leaf area index. Tagesson et al. (2016b) performed a quality check of the EC data due to the high net CO₂ exchange measured at the Dahra field site and explained the high values by a combination of moderately dense herbaceous C4 ground vegetation, high soil nutrient availability and a grazing pressure resulting in compensatory growth and fertilization effects. Another possible explanation could be that the West African monsoon brings a humid layer of surface air from the Atlantic, possibly increasing vegetation production for the most western part of the Sahel (Tagesson et al., 2016a).

Our model substantially overestimated GPP for Wankama Millet (Fig. 7f). Being a crop field, this site differed from the other sites in its species composition and ecosystem structure, as well as land and vegetation management. Crop fields in southwestern Niger are generally characterized by rather low production, resulting from decreased fertility and soil loss caused by intensive land use (Cappelaere et al., 2009). These specifics of the Wankama Millet site may cause the model, parameterized with observations from the other study sites without this strong anthropogenic influence, to overestimate GPP at this site. Similar results were found by Boulain et al. (2009) when applying an upscaling model using leaf area index for Wankama Millet and Wankama Fallow. It worked well for Wankama fallow, whereas it was less conclusive for Wankama Millet. The main explanation for this difference was low leaf area index in millet fields because of a low density of millet stands due to agricultural practice. There is extensive savanna clearing for food production in the Sahel (Leblanc et al., 2008; Boulain et al., 2009; Cappelaere et al., 2009). To further understand impacts of this land cover change on vegetation production and land–atmosphere exchange processes, there is an urgent need for more study sites covering cropped areas in this region.

In Demokeya, GPP was slightly underestimated for 2008 (Fig. 7c) because modelled F_{opt} was much lower than the actual measured value in 2008 (the thick black line in Fig. 4). An improvement of the model could be to incorporate some parameters that constrain or enhance F_{opt} depending on environmental stress. Indeed, the regression tree analysis indicated that incorporating hydrometeorological variables increased the ability to predict both F_{opt} and α. Conversely, for spatial upscaling purposes, it has been shown that including modelled hydrometeorological constraints on LUE decreases the ability to predict vegetation production due to the incorporated uncertainty in these modelled variables (Fensholt et al., 2006; Ma et al., 2014). For spatial upscaling to regional scales, it is therefore better to simply use relationships with EO data. This is particularly the case for the Sahel, one of the
The pattern seen in the spatially explicit GPP budgets (Fig. 5c) may be influenced by a range of biophysical and anthropogenic factors. The clear north–south gradient is expected given the strong north–south rainfall gradient in the Sahel. The West African monsoon mentioned above could also be an explanation of high GPP values in the western part of the Sahel, where values were relatively high in relation to GPP at similar latitudes in the central and eastern Sahel (Fig. 5c). The areas with the highest GPP are sparsely populated woodlands or shrubby savanna with a relatively dense tree cover (Brandt et al., 2016). However, the maps produced here should be used with caution as they are based on upscaling of data collected at only six EC sites available in the region, especially given the issues related to the cropped fields discussed earlier. Still, the average GPP budget for the entire Sahel 2001–2014 was close to an average annual GPP budget estimated at these six sites (692 ± 89 g C m⁻² yr⁻¹; Tagesson et al., 2016a). The range of GPP budgets in Fig. 5c is also similar to previous annual GPP budgets reported from other savannas across the world (Veenendaal et al., 2004; Chen et al., 2003, 2006; Kanniah et al., 2010).

Although MOD17A2 GPP has previously been shown to capture GPP in several ecosystems types well (Turner et al., 2006, 2005; Heinsch et al., 2006; Sims et al., 2006; Kanniah et al., 2009), it has been shown to underestimate it in others (Coops et al., 2007; Gebremichael and Barros, 2006; Sjöström et al., 2013). Gross primary production of Sahelian drylands have not been captured well by MOD17A2 (Sjöström et al., 2013; Fensholt et al., 2006), and as we have shown, this underestimation persists in the latest MOD17A2H GPP (collection 6) product (Fig. 2). The main reason for this pronounced underestimation is that maximum LUE is set to 0.84 g C MJ⁻¹ (open shrubland, Demokekeya) and 0.86 g C MJ⁻¹ (grassland; Agoufou, Dahra, Kelma; Wankama Millet and Wankama Fallow) in the BPLUT, i.e. much lower than maximum LUE measured at the Sahelian measurement sites of this study (average: 2.47 g C MJ⁻¹; range: 1.58–3.50 g C MJ⁻¹; Sjöström et al., 2013; Tagesson et al., 2015a), a global estimate of ~1.5 g C MJ⁻¹ (Garbulsky et al., 2010) and a savanna site in Australia (1.26 g C MJ⁻¹; Kanniah et al., 2009).

Several dynamic global vegetation models have been used for decades to quantify GPP at different spatial and temporal scales (Dickinson, 1983; Sellers et al., 1997). These models are generally based on the photosynthesis model of Farquhar et al. (1980), a model particularly sensitive to uncertainty in photosynthetic capacity (Zhang et al., 2014). This and several previous studies have shown that both photosynthetic capacity and efficiency (both α and LUE) can vary considerably between seasons as well as spatially vary, and both within and between vegetation types (Eamus et al., 2013; Garbulsky et al., 2010; Ma et al., 2014; Tagesson et al., 2015a). This variability is difficult to estimate using broad values based on land cover classes, yet most models apply a constant value, which can cause substantial inaccuracies in the estimates of seasonal and spatial variability in GPP. This is particularly a problem in savannas that consist of several plant functional types (C3 and C4 species, and a large variability in tree and/or herbaceous vegetation fractions; Schles and Archer, 1997). This study indicates the applicability of EO as a tool for parameterizing spatially explicit estimates of plant physiological variables, which could improve our ability to simulate GPP. Spatially explicit estimates of GPP at a high temporal and spatial resolution are essential for environmental change studies in the Sahel and can contribute to increased knowledge regarding changes in GPP, its relationship to climatic change and anthropogenic forcing, and simulations of ecosystem processes and hydro-biochemical cycles.

Data availability. The EC data are available from Fluxnet (http://fluxnet.ornl.gov) and CarboAfrica (http://www.carboafrica.net/indexen.asp). The ERA Interim PAR data are available from ECMWF (http://apps.ecmwf.int/datasets/data/interim-full-daily/).
The data products from this study are available from the corresponding author upon request.

The Supplement related to this article is available online at doi:10.5194/bg-14-1333-2017-supplement.

Competing interests. The authors declare that they have no conflict of interest.

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ECMWF: ERA-Interim: surface photosynthetically active radiation (surface PAR) values are too low, available at: https://software.ecmwf.int/wiki/display/CKB, last access: 7 November 2016b.


